

Science of the Total Environment

Surface and groundwater flow exchanges and lateral hydrological connectivity in environments of the Matusagaratí Wetland, Panama --Manuscript Draft--

Manuscript Number:	
Article Type:	Research Paper
Keywords:	Tropical wetland; Hydrodynamics; Hydrochemistry; Stable Isotopes; Ecohydrology
Corresponding Author:	Eleonora Carol, Ph Geological Research Center La Plata, ARGENTINA
First Author:	Eleonora Carol, Ph
Order of Authors:	Eleonora Carol, Ph María del Pilar Alvarez Manuel Arcia Indra Candanedo
Abstract:	<p>The Matusagaratí wetland in the Panamanian Darien is one of the largest wetlands in Central America. These types of riverine wetlands, associated with large drainage basins, are complex hydrological environments where variations in water flows and exchanges condition the existence of different wetland habitats. The aim of the work was to establish the hydrological functioning of the Matusagaratí wetland in different sectors of the Balsas River, emphasizing the exchanges of surface and groundwater flows and the hydrological connectivity that exists between the different laterally linked wetland environments. For this purpose, a monitoring network for surface water and groundwater was established along transects intersecting various wetland environments in the middle and lower basin of the Balsas River. This network is complemented by measurement points for surface water located in streams and in the upper basin of the river.</p> <p>Data collected in sensors installed in boreholes were compared to river level and precipitation data. Continuous water level recording sensors were installed at the monitoring points, and samples were collected for the determination of major ions and stable isotopes. The results indicate that in the mangroves of the lower basin and in the cativo forests of the middle basin levee, there is a strong exchange of water between the river and the shallow groundwater. Meanwhile, in the middle basin, mixed forests and orey forests developed on the alluvial plain exhibit a hydrological functioning that depends primarily on precipitation inputs. This study provides data that could serve as a basis for the management of this large tropical wetland, which, despite having protection initiatives, could be hydrologically impacted by the unsustainable socio-economic practices.</p>
Suggested Reviewers:	Luigi Tosi Institute of Geosciences and Earth Resources National Research Council Padua Branch luigi.tosi@igg.cnr.it
	Sandra Donnici Institute of Geosciences and Earth Resources National Research Council Padua Branch sandra.donnici@igg.cnr.it
	Lucia Santucci National Scientific and Technical Research Council luciasantucci@fcnym.unlp.edu.ar
	Carolina Tanjal National Scientific and Technical Research Council ctanj@cig.museo.unlp.edu.ar
	Cristina Dapeña

	University of Buenos Aires dapenna@gmail.com
	Eduardo Mariño National University of La Pampa emarinio@exactas.unlpam.edu.ar
Opposed Reviewers:	



Consejo Nacional de Investigaciones Científicas y Técnicas

December 13, 2023

Dear Editor,

We submit the manuscript entitled "Surface and groundwater flow exchanges and lateral hydrological connectivity in environments of the Matusagaratí Wetland, Panama" by Carol E., Alvarez M.P., Arcia M. and Candanedo I., to be considered for publication in Science of The Total Environment.

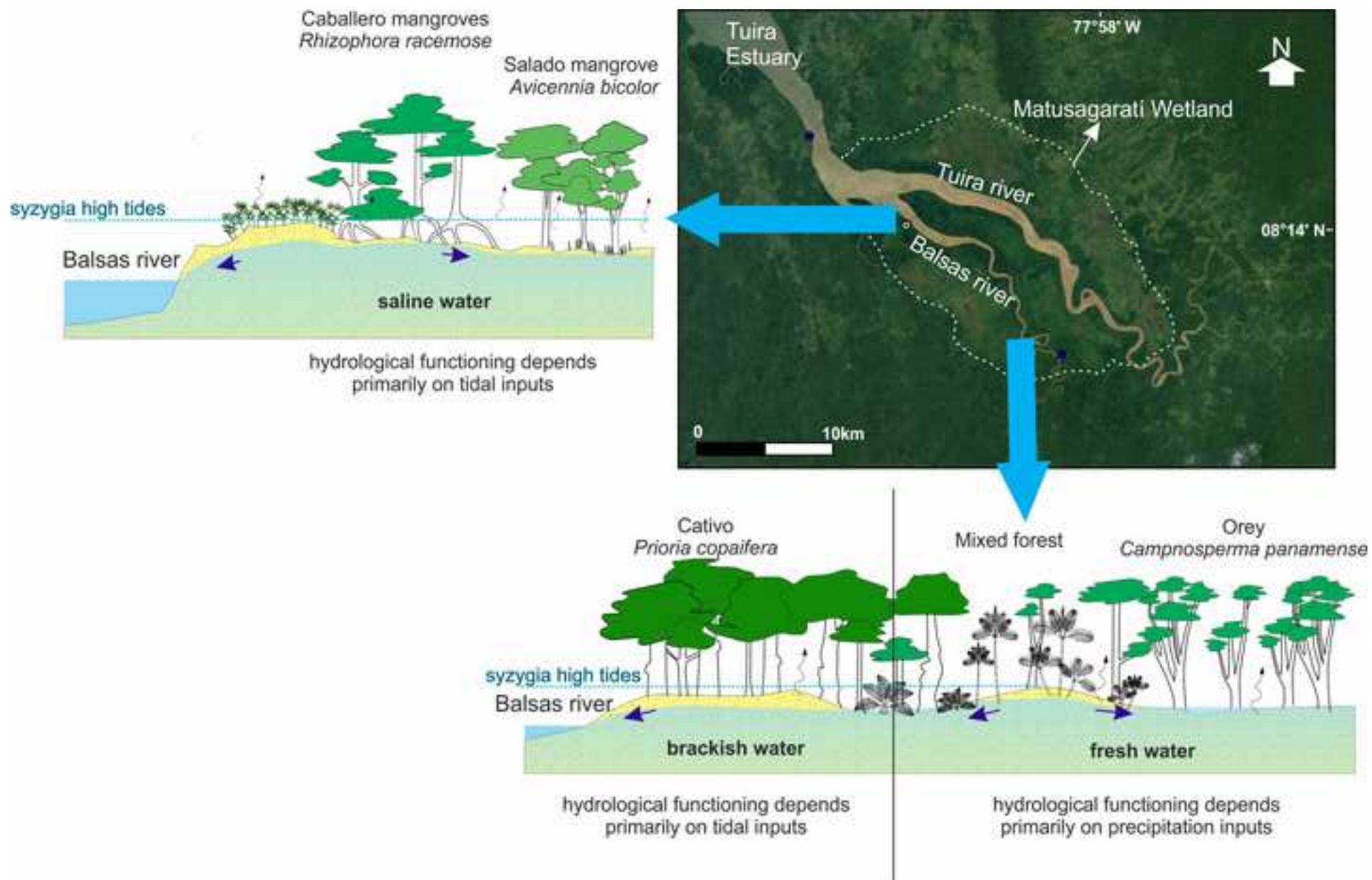
The work establishes the hydrological functioning of the Matusagaratí wetland in different sectors of the Balsas River, emphasizing the exchanges of surface and groundwater flows and the hydrological connectivity that exists between the different laterally linked wetland environments. The Matusagarati wetland in the Panamanian Darien is one of the largest wetlands in Central America. The wetland area is developed in the Darien jungle, which is very difficult to access, a situation that limits the carrying out of research work.

In our study, surface and groundwater monitoring networks were generated, where water level sensors were installed and the majority and isotopic chemistry of the water were analyzed. The joint analysis of the hydrological behavior with the wetland vegetation allowed us to recognize environments with different water inputs. The results obtained not only provide data in unstudied areas but also contribute to generating water management guidelines in an area where, although there are natural reserve areas, there are also anthropogenic pressures on land use.

This new article whose content has not been published, may be of particular interest to the readers of the journal.

Sincerely,

Carol Eleonora
Corresponding author



Highlight

- Matusagaratí wetland is one of the largest wetlands in Central America.
- Different wetland environments are recognized in the Balsas River.
- The vegetation of the wetland is conditioned by the different water inputs.
- The water that contributes to the wetland in the lower basin is mainly of tidal origin.
- In the middle basin the wetland is mainly dependent on river and rainfall input.

[Click here to view linked References](#)

1 **Surface and groundwater flow exchanges and lateral hydrological connectivity in**
2 **environments of the Matusagaratí Wetland, Panama**

3

4 Eleonora Carol^{1*}, María del Pilar Alvarez², Manuel Arcia³, Indra Candanedo³

5

6 1- Centro de Investigaciones Geológicas (CIG). Consejo Nacional de Investigaciones Científicas y
7 Técnicas (CONICET), Universidad Nacional de La Plata (UNLP), Argentina.

8

9 2- Instituto Patagónico para el Estudio de los Ecosistemas Continentales (IPEEC), Consejo
10 Nacional de Investigaciones Científicas y Técnicas (CONICET), Universidad Nacional de la
11 Patagonia San Juan Bosco (UNPSJB), Argentina.

11

3- Technological University of Panama (UTP), Republic of Panama.

12

*Corresponding author: eleocarol@fcnym.unlp.edu.ar

13

Abstract

14

The Matusagarati wetland in the Panamanian Darien is one of the largest wetlands in Central
15 America. These types of riverine wetlands, associated with large drainage basins, are complex
16 hydrological environments where variations in water flows and exchanges condition the existence
17 of different wetland habitats. The aim of the work was to establish the hydrological functioning
18 of the Matusagaratí wetland in different sectors of the Balsas River, emphasizing the exchanges
19 of surface and groundwater flows and the hydrological connectivity that exists between the
20 different laterally linked wetland environments. For this purpose, a monitoring network for
21 surface water and groundwater was established along transects intersecting various wetland
22 environments in the middle and lower basin of the Balsas River. This network is complemented
23 by measurement points for surface water located in streams and in the upper basin of the river.

24

Data collected in sensors installed in boreholes were compared to river level and precipitation
25 data. Continuous water level recording sensors were installed at the monitoring points, and
26 samples were collected for the determination of major ions and stable isotopes. The results
27 indicate that in the mangroves of the lower basin and in the cativo forests of the middle basin

28 levee, there is a strong exchange of water between the river and the shallow groundwater.
29 Meanwhile, in the middle basin, mixed forests and orey forests developed on the alluvial plain
30 exhibit a hydrological functioning that depends primarily on precipitation inputs. This study
31 provides data that could serve as a basis for the management of this large tropical wetland, which,
32 despite having protection initiatives, could be hydrologically impacted by the unsustainable socio-
33 economic practices.

34 **Keywords:** Tropical wetland; Hydrodynamics; Hydrochemistry; Stable Isotopes; Ecohydrology.

35

36 **1. Introduction**

37 Wetlands are critical environments that provide ecosystem services valued at trillions of dollars
38 annually (De Groot et al., 2009; Barbier, 2011; Ivory et al., 2019). They help mitigate flood risk,
39 provide key freshwater resources, play an essential role in nutrient and carbon cycling, and
40 support many local and regional economies (Costanza et al., 1997; Erwin 2009; Ahmed, 2015;
41 Adame et al., 2019; Cuthbert et al., 2022). Therefore, the degradation and loss of wetlands not
42 only causes the deterioration and loss of biodiversity, but also results in the loss of associated
43 ecosystem services, with economic implications (Wasserman and Dalu, 2022).

44 In river wetlands, water exchanges between the river and the shallow groundwater that sustain
45 ecosystems are key processes that determine not only the environmental characteristics of wetland
46 systems, but also the connectivity among different habitats (Dube et al., 2019). In this sense,
47 studying the hydrological functioning of wetlands considering surface and groundwater flows is
48 valuable not only to define their functioning, but also to predict the impact that natural or human-
49 induced hydrological modifications can cause in habitats that support wetlands. The sustainable
50 use of wetlands and their management requires planning and a good scientific understanding
51 about how these ecosystems function (Ahmed, 2015).

52 The Matusagarati wetland in the Panamanian Darien is one of the largest wetlands in Central
53 America. In fact, Matusagarati is a complex of wetlands that includes several types of flooded

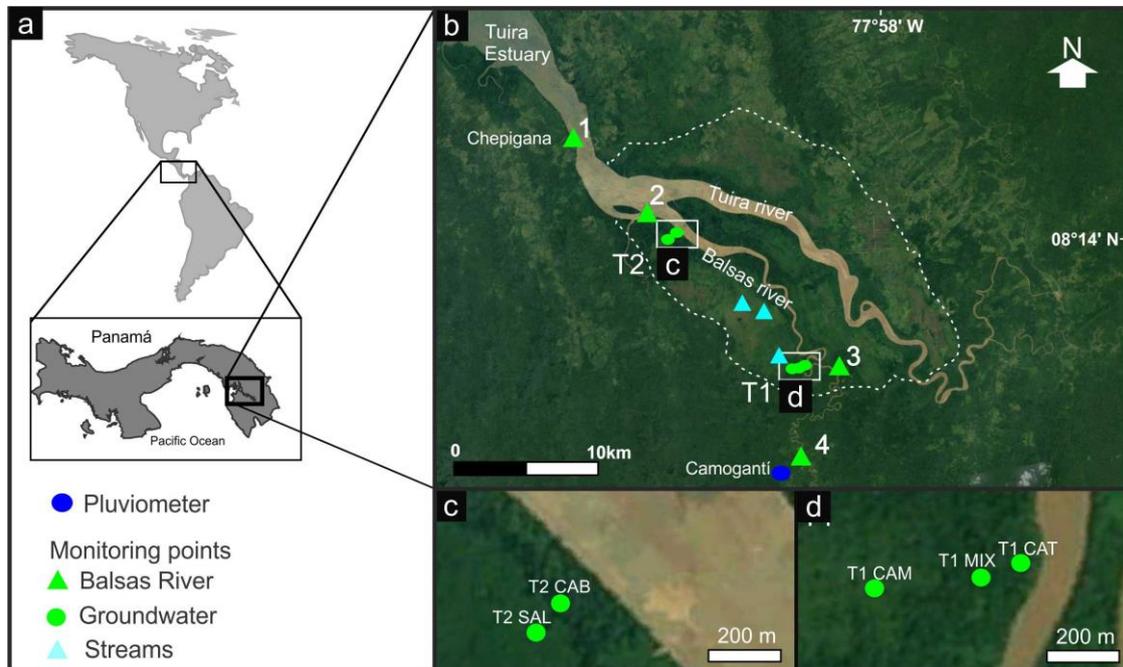
54 forests, flooded grasslands and mangroves covering approximately 56,250 hectares (Candanedo
55 2021). Matusagarati lies along the margins and adjacent lowlands of the Tuirá and Balsas rivers
56 before emptying into the Pacific Ocean, forming an extensive estuary where fresh and marine
57 waters meet. More inland, this wetland also receives freshwater from small streams draining from
58 nearby ridges.

59 There are three protected areas that provide legal protection to 70% of the wetland: the Filo del
60 Tallo-Canglon hydrological reserve, the Chepigana Forest Reserve, and the Matusagarati Wildlife
61 Refuge. Despite this legal protection, there is no on-the-ground management resulting in
62 inappropriate practices such burning of herbaceous wetlands during the dry months by local cattle
63 ranchers who seek to expand their cattle pastures. Furthermore, during the last decade, nearly a
64 third of the wetland has been illegally titled and part of it drained and transformed into rice
65 cultivation (Ministerio de Ambiente 2016; CREHO 2015), dramatically changing its hydrological
66 behavior.

67 Until recently, there has been a lack of baseline studies to support better management and
68 involvement of local communities in sustainable management due to the lack of roads which make
69 access to study sites extremely difficult and costly. However, five years ago, a group of national
70 as well as international researchers have joined efforts to understand Matusagarati's biodiversity
71 and ecological functioning. Studies have been carried out on plants (Grauel and Putz 2004; Ortiz
72 et al. 2020; Ortiz et al. 2022; Ibáñez and Flores 2020), bird communities (Aparicio 2020), and the
73 wetland's potential as carbon storage and fisheries (López and Cunampio 2023). A map on the
74 wetland's vegetation types has also been developed.

75 The only previous study on hydrological functioning in Matusagarati was conducted along the
76 Tuirá River (Carol et al., 2020; 2022). These studies show that some wetland environments
77 depend upon the exchanges of surface water - groundwater flows while other wetland types were
78 dependant mainly on precipitation. Though the hydrological dynamics of the Tuirá River's
79 wetlands has started to be understood, the Balsas River wetlands system remain to be studied.
80 The aim of the work was to establish the hydrological functioning of the Matusagarati wetland in

81 different sectors of the Balsas River, emphasizing the exchanges of surface and groundwater flows
82 and the hydrological connectivity that exists between the different laterally linked wetland
83 environments.



84
85 Figure 1: (a) Location of the study area. (b) Location of the sampling points 1: Balsas River near
86 estuary, 2: Balsas River lower basin, 3: Balsas River middle basin, 4: Balsas River upper basin,
87 T1: Transect 1 and T2: Transect 2. The white dotted line indicates the location of the Matusagaratí
88 wetland. (c) Details of transect 2. (d) Details of transect 1.

89

90 2. Methodology

91 To analyze the hydrodynamics and hydrochemistry of the wetland, two transects were defined in
92 different sectors of the wetland. These transects were arranged perpendicular to the bed of the
93 Balsas River and intercepting different environments within the wetland. Transect 1 intercepts
94 three wetland environments. The river levee, the raised strip that develops adjacent to the river
95 margin, is covered by an almost monotypic cativo forest (*Prioria copaifera*) with some
96 alcornoque trees (*Mora oleifera*), and palms such as *Astrocaryum standleyanum*, *Euterpe*
97 *oleracea* y *Elaeis oleifera*. Behind the cativo forest there is a flooded evergreen mixed forest that

98 shows a wider variety of species. These include sangrillo (*Pterocarpus officinalis*), alcornoque
99 (*Mora oleifera*), cuchillito (*Pentaclethra maculosa*) and tangaré (*Carapa guianensis*). Finally,
100 located in a depression behind the flooded evergreen mixed forest, the ore forest
101 (*Camposperma panamense*) is found. This type of forest has just been recently mapped. The
102 ore trees grow together with palm species such as *Euterpe oleracea* and occasionally *Elaeis*
103 *oleifera*.

104 Transect 2 down the river, closer to the estuary, starts with a stand of mangrove fern (*Acrostichum*
105 *danaeifolium*), followed by 20-25 m high caballero mangroves (*Rhizophora racemosa*) with their
106 characteristic aerial roots that grow from the main stem, resembling flying buttresses. More inland
107 there is a nearly monotypic salado mangrove (*Avicennia bicolor*) between 15 and 20 m high with
108 pneumatophores, specialized aerial root-like structures that allow the absorption of gases directly
109 from the atmosphere, an adaptation to flooded environments.

110 In each transect, wells were installed within each of the wetland environments described above.
111 The boreholes were made between 3 and 4 m deep with a manual auger. The boreholes were cased
112 with a continuous grooved filter PVC pipe wrapped in a fine plastic mesh, filling the annular
113 space with a pre-filter of siliceous gravel, and sealing the upper part of the annular space of the
114 well. Hobbo level data logger and Odyssey® Capacitance Water Level sensors that continuously
115 register studied variables were installed in all the boreholes, which were programmed to measure
116 the water level every 15 minutes. To measure the Balsas River levels, similar sensors were
117 installed up the river in transect 1 and down the river in transect 2 (Fig. 1).

118 Precipitation was measured using an Odyssey® tipping bucket rain gauge classic, located in
119 Camogantí, in the upper basin sector of the Balsas River. Data collected in the sensors installed
120 in the boreholes were later compared to the river level and precipitation data. The comparison
121 between river and groundwater levels allowed the analysis of flow exchanges between surface
122 water and groundwater. Comparison of groundwater levels with rainfall data was carried out to
123 visualize the influence of rain infiltration.

124 Water samples were taken at 3 points of the Balsas River located in the upper, middle, and lower
125 basin, and in small streams that drain from the wetland to the main river, and in the wells

126 (groundwater) installed in the transects. At all points electrical conductivity (EC) was measured
127 *in situ* using a portable conductivity meter instrument (Lutron® WA-2017SD). Major ions and
128 stable isotopes of the water molecule were determined in the water samples in the laboratory. This
129 work uses data from water samples collected in June 2022, June 2023, and August 2023. The
130 determination of major ions was carried out using standardized methods (APHA, 1998) in the
131 Geochemistry Laboratory of the Geological Research Center in La Plata, Argentina. Stable
132 isotopes of the water molecule were determined by mass spectroscopy in the Stable Isotope
133 Laboratory of the University of San Luis (Argentina), using a Thermo Finnigan MAT Delta Plus
134 XL continuous flow mass spectrometer. The analytical accuracy is $\pm 0.05\text{‰}$ and $\pm 0.5\text{‰}$, for $\delta^{18}\text{O}$
135 and $\delta^2\text{H}$, respectively. Isotopic results are presented as $\delta\text{‰}$, defined as $\delta=1000(R_s-R_r)/R_r\text{‰}$,
136 where δ is the isotopic deviation in ‰ relative to Vienna Standard Mean Ocean Water (V-SMOW)
137 (Gonfiantini, 1978); R is the isotopic ratio ($^2\text{H}/^1\text{H}$, $^{18}\text{O}/^{16}\text{O}$); r: international reference and s:
138 sample. The isotopic values were compared with the local meteoric line $\delta^2\text{H} = 7.63 \delta^{18}\text{O} + 6.51$
139 for the Pacific coast of central Panamá (Kern et al., 2016).

140

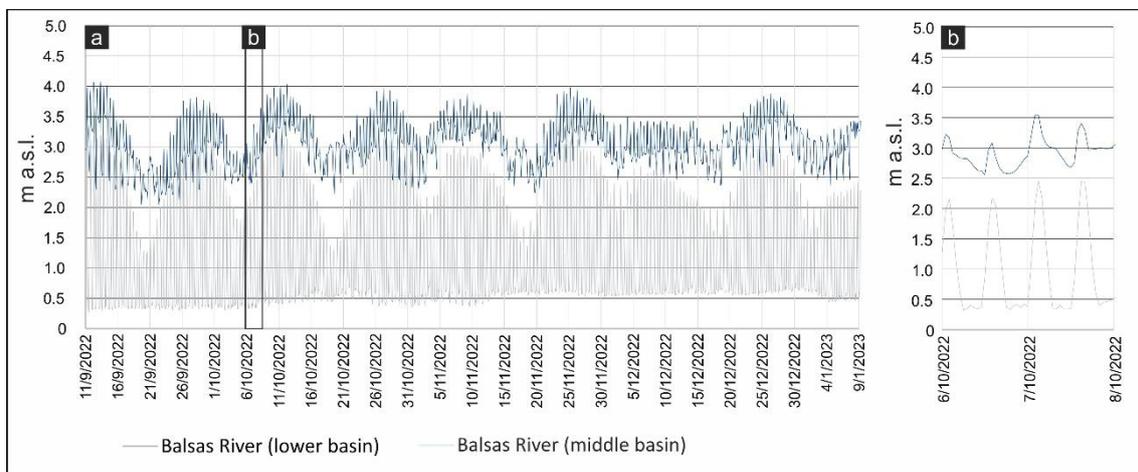
141 **3. Results**

142 *3.1. Wetlands hydrodynamic*

143 The description of the hydrodynamic results will begin with the analysis of the variations in the
144 values of water levels in the river, comparing data from the middle basin and the lower middle
145 basin where continuous recording sensors were installed. Subsequently, the relationship between
146 river water and groundwater in the different wetland environments intercepted by transect 1 (in
147 the lower basin) and transect 2 (in the middle basin) will be analyzed. For these last cases, it will
148 also be analyzed whether there is any relationship between water levels and precipitation.

149 The records of water levels in the Balsas River (Fig. 1a) show oscillations related to tidal
150 influence both in the lower basin sector and in the middle basin sector close to transect 1. The
151 relative position of water levels in the river between both sectors shows that the upstream sensor
152 generally registers a higher water level than the sensor in the downstream sector. This would

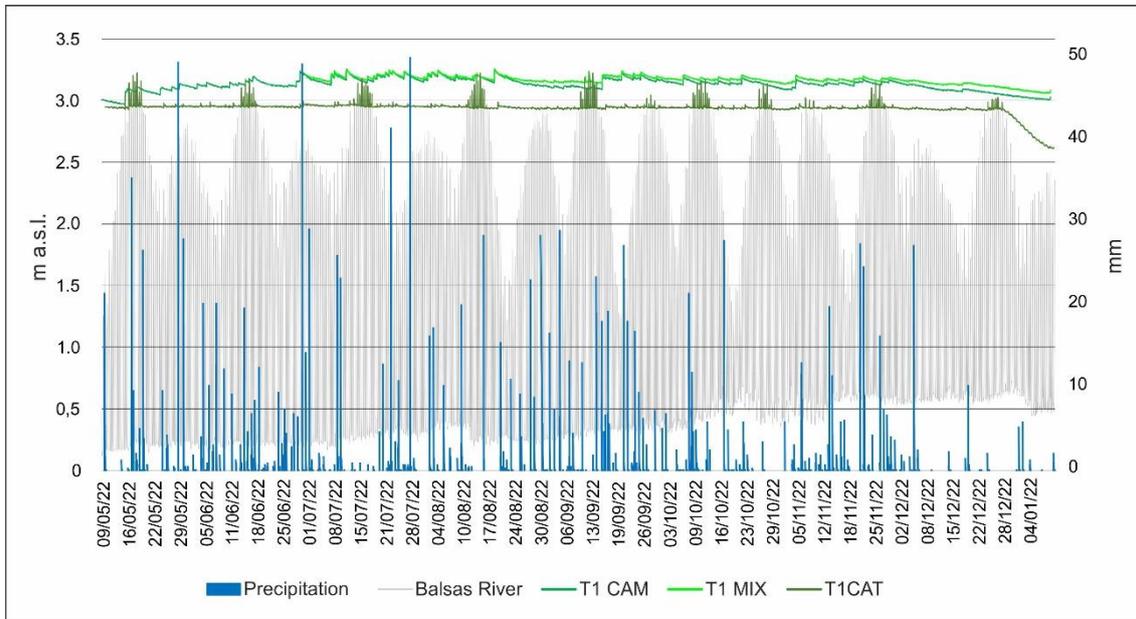
153 indicate that the propagation of the tide along the river channel do not allow the river water to
 154 drain but accumulate upstream, causing rises in river levels in the middle section of the basin. In
 155 the lower basin, the rise in level is due to the entry of tidal water, while towards the middle - upper
 156 basin the rise is due to the obstruction of the drainage by the entry of the tide. In both sectors of
 157 the basin, the high tide and low tide peaks occur at the same time (Fig 2b). This supports the idea
 158 of the above-mentioned behavior and not that of a rise caused only by tidal propagation within
 159 the channel, which would be observed with a phase shift in the peaks caused by the delay time in
 160 which the tide takes to propagate within the channel.



161
 162 Figure 2: Variations of the water level in the Balsas River (in m a.s.l.) in sectors of the lower basin
 163 and middle basin.

164
 165 The comparison of the water levels in the Balsas River with the phreatic levels in the wells of
 166 transect 1 show that there are different behaviors (Fig. 3). The sensor located on the river levee in
 167 an environment dominated by cativo forest (T1 CAT) shows tide-related rises at all high tides that
 168 exceed 2.75 m a.s.l. At low tide or high tides below said level, the water table is mostly at values
 169 close to 2.90 m a.s.l., with the depth of the groundwater close to 0.08 m. Particularly starting in
 170 January 2023, when high tides are low and rainfall is very scarce, a decrease in the groundwater
 171 level in the cativo forest is observed. Minewhile, the groundwater in the mixed forest (T1 MIX)
 172 and orey forest (T1 CAM) has a very shallow water table (outcropping sometimes). In these

173 environments, the water table does not present oscillations in relation to the tidal flow, but the
 174 rises can be associated in most cases with precipitation. An exception to this behavior was
 175 observed on the first half of September 2022, when the high tides reached the highest point on
 176 record and generated small rises in the water table in the mixed forest sensor.



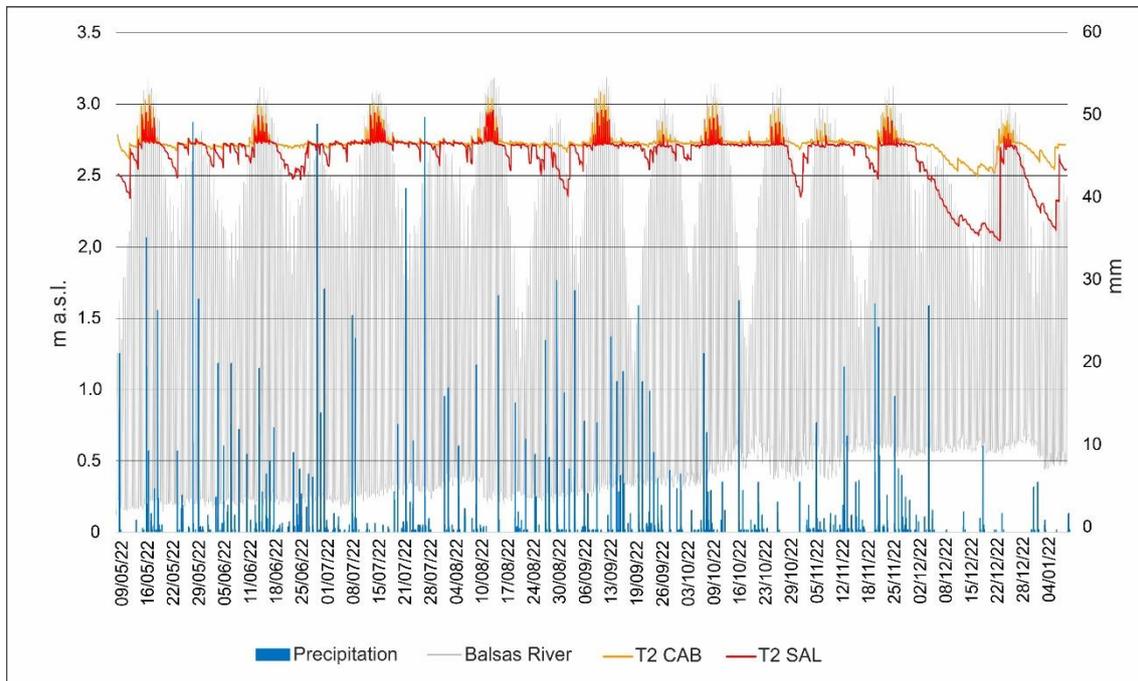
177

178 Figure 3: Variations of the groundwater level in different environments of transect 1, the Balsas
 179 River level (in m a.s.l.), and daily precipitation data (in mm).

180

181 In transect 2, the comparison of the phreatic levels with the river levels shows a different behavior
 182 (Fig. 4). The groundwater level in the caballero mangrove (T2 CAB) is less than 0.05 m deep and
 183 in the salado mangrove (T2 SAL) it is less than 0.22 m deep. In both environments, the water
 184 table rises when the river level exceeds 2.75 m a.s.l. during high tide (mainly syzygy tide). At low
 185 tide and high tides less than 2.75 m a.s.l., the groundwater had a level close to 2.70 m a.s.l.,
 186 between May and November 2022. In this section of the record (Fig. 4) the syzygy high tides
 187 exceed 2.75 m a.s.l. and rainfall is abundant. Starting in December 2022, a change in the
 188 groundwater level is observed, which tends to decrease mainly in the salado mangrove. This
 189 second section of the record is characterized by lower syzygy high tides and for being a drier

190 period with rainfall that rarely exceeds 10 mm. However, the water table rises rapidly when high
191 tides above 2.75 m a.s.l. occur within this period (see date of January 20, 2023).



192

193 Figure 4: Variations of the groundwater level in transect 1 and the level of the Balsas River (in m
194 a.s.l.) and daily precipitation data (in mm).

195

196 3.2.3.2. Hydrochemistry and stable isotopes of the water molecule

197 In the water of the Balsas River, an increase in the EC of the water is recorded from the upper
198 basin to the lower basin. In the upper basin, the EC varies between 115 and 125 $\mu\text{S}/\text{cm}$, in the
199 middle basin between 118 and 484 $\mu\text{S}/\text{cm}$ and in the lower basin between 13940 and 21200 $\mu\text{S}/\text{cm}$.
200 The small streams that drain towards the river have low EC with values between 69 and 282
201 $\mu\text{S}/\text{cm}$. These differences in the EC are reflected in the contents of major ions and hydrochemical
202 facies (Fig. 5).

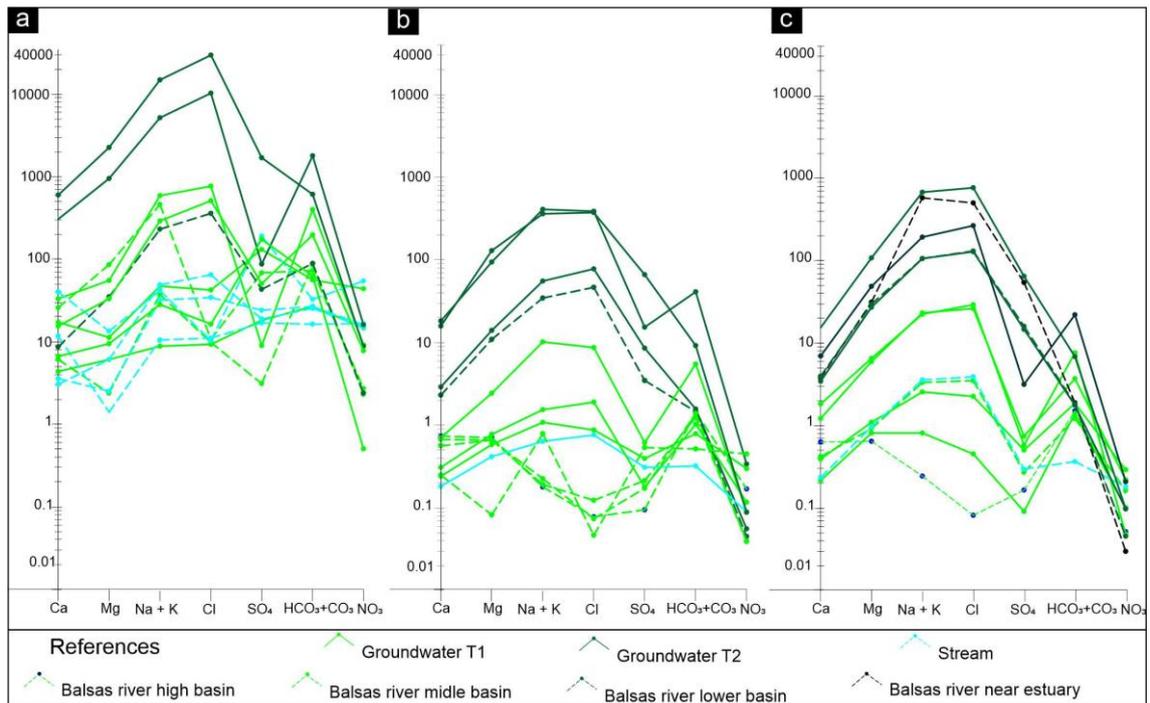
203 In the Balsas River, the dominant ions in the upper basin vary, depending on the sampling,
204 between Ca^{+2} (0.64 - 0.73 meq/L) and Mg^{+2} (0.64 - 0.70 meq/L) and HCO_3^- (1.43 - 1.51 meq/L),
205 or between Na^+ and Cl^- (0.16 – 0.22 and 0.07 – 0.08 meq/L). While, in the middle and lower basin

206 an increase in the concentrations of Na^+ (80.22 and 103.18 meq/L in the middle and lower basin
207 respectively) and Cl^- (4.04 and 153.96 meq/L in the middle and lower basin respectively) is
208 observed, being these markedly dominant over the other major ions, particularly in the lower
209 basin. These spatial variations determine a change in hydrochemical facies in the river water with
210 sodium-calcic bicarbonate and sodium chloride trends in the upper basin that transition to sodium
211 chloride in the middle and lower basin. In the streams that flow to the main river, dominant ions
212 are sodium chloride to sodium bicarbonate-chloride, with the average content of 22.32 meq/L for
213 Na^+ , 16.95 meq/L for HCO_3^- and 20.86 meq/L for Cl^- .

214 The groundwater in transect 1 presents different EC values in each wetland environment. In three
215 samples of groundwater taken in the cativo forest (T1 CAT), higher EC was recorded than in the
216 groundwater in the mixed forest (T1 MIX) and orey forest (T1 CAM). The EC in the cativo forest
217 varied between 1423 and 3680 $\mu\text{S}/\text{cm}$, while in the mixed and orey forest the electrical
218 conductivity was similar in May 2022 and June 2023 with values between 55 and 311 $\mu\text{S}/\text{cm}$.
219 However, in August 2023 a significant increase was recorded in the mixed forest reaching values
220 of 2600 $\mu\text{S}/\text{cm}$. The hydrochemical facies in relation to the content of major ions are sodium
221 chloride in the cativo forest and sodium bicarbonate-chloride in the mixed and orey forest (Fig.
222 5).

223 In transect 2, the groundwater EC values are up to two orders of magnitude higher than in transect
224 1. In the caballero mangrove environment (T2 CAB) the EC varied between 14142 and 34250
225 $\mu\text{S}/\text{cm}$, while in the salado mangrove (T2 SAL) between 29870 and 46200 $\mu\text{S}/\text{cm}$. In both cases,
226 the high EC is associated with a high concentration of Na^+ ions (mean values of 1894.29 meq/L
227 in T2 CAB and 5200.15 meq/L T2 SAL) and Cl^- (mean values of 3654.34 meq/L in T2 CAB and
228 of 10352.76 meq/L T2 SAL). In this sector of the wetland there is a dominance of sodium chloride
229 facies in both the river and the groundwater (Fig. 5).

230



231

232 Figure 5: Schoeller plots representing the concentration of major ions in meq/L and the designs
 233 of the hydrochemical facies. (a) May 2022, (b) June 2023, (c) August 2023.

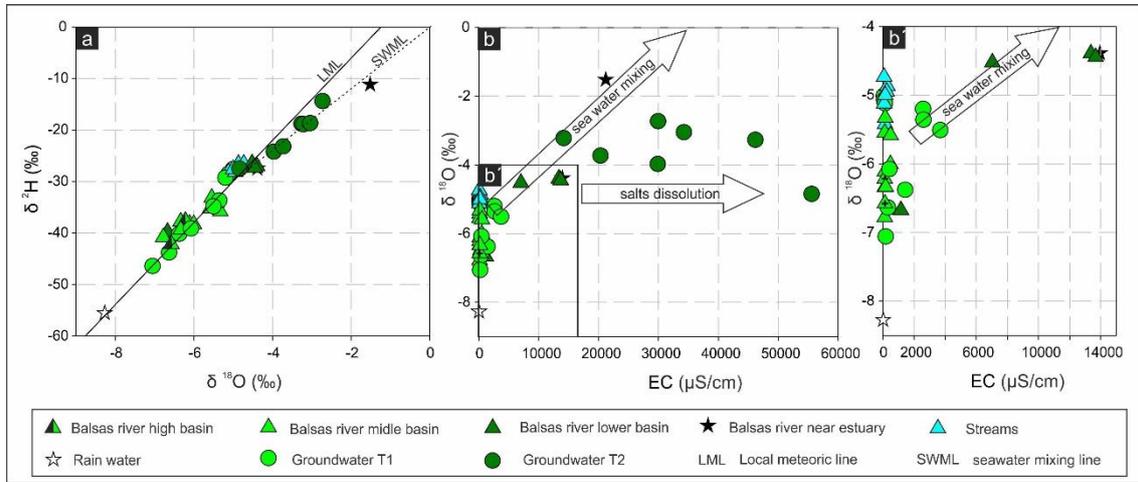
234

235 The $\delta^{18}\text{O}$ vs. $\delta^2\text{H}$ graph (Fig. 6a) shows the results of the isotopic signal for samples taken in the
 236 upper and middle basin of the Balsas River, in the streams and in the groundwater samples of
 237 transect 1. All of them are located around the local meteoric line. The rainwater sampled in May
 238 2022 is also located on this line, however, it shows the least enriched values.

239 The isotopic signal of the samples from the lower basin of the Balsas River and the groundwater
 240 of transect 2, on the other hand, deviates from the trend described above as they are located around
 241 a mixing line between river water and sea water (Fig. 6a). In contrast, the graph of $\delta^{18}\text{O}$ vs. EC
 242 (Fig. 6b) shows that the samples from the upper and middle basin of the Balsas River, as well as
 243 those from the streams and the groundwater in transect 1, present isotopic variations associated
 244 with a small increase in the EC.

245 The groundwater samples from transect 1, located in the cativo forest (Fig. 6b'), slightly deviate
 246 from this general trend, and tend towards mixing with seawater. Likewise, all groundwater

247 samples from transect 2 present strong increases in EC associated with few variations in $\delta^{18}\text{O}$
 248 according to a trend of salt dissolution (Fig. 6b).



249
 250 Figure 6: (a) $\delta^2\text{H}$ vs. $\delta^{18}\text{O}$, (b) $\delta^{18}\text{O}$ vs. EC, (b') detail of figure (b).

251

252 **4. Discussion**

253 The hydrological dynamics of wetlands associated to estuaries and large river basins is very
 254 complex (Ivory et al., 2019, Xiao et al., 2019). Because in wetlands hydrology plays a key role in
 255 the development of vegetation, it is possible to recognize different wetland environments with
 256 their own hydrological behaviors and plant communities (Chui et al., 2011). In this sense, the
 257 results show that the Matusagaratí wetland presents different environments where different
 258 degrees of interaction between surface and groundwater can be recognized with variations in
 259 lateral hydrological connectivity.

260 Although high rainfall, typical of the humid tropical climate with values close to 2900 mm per
 261 year (Fabrega et al., 2013), determines much of Matusagaratí's hydrology, tidal influence is also
 262 significant (Carol et al., 2022) as shown by the hydrodynamic and hydrochemical differences
 263 recorded in different sectors of the Balsas River. In terms of hydrodynamics, the river in lower
 264 and middle basin showed oscillations that respond to the entry of the tide from the Pacific Ocean
 265 (Fig. 2). The entry of tidal water through the river channel produces an increase in salinity in the

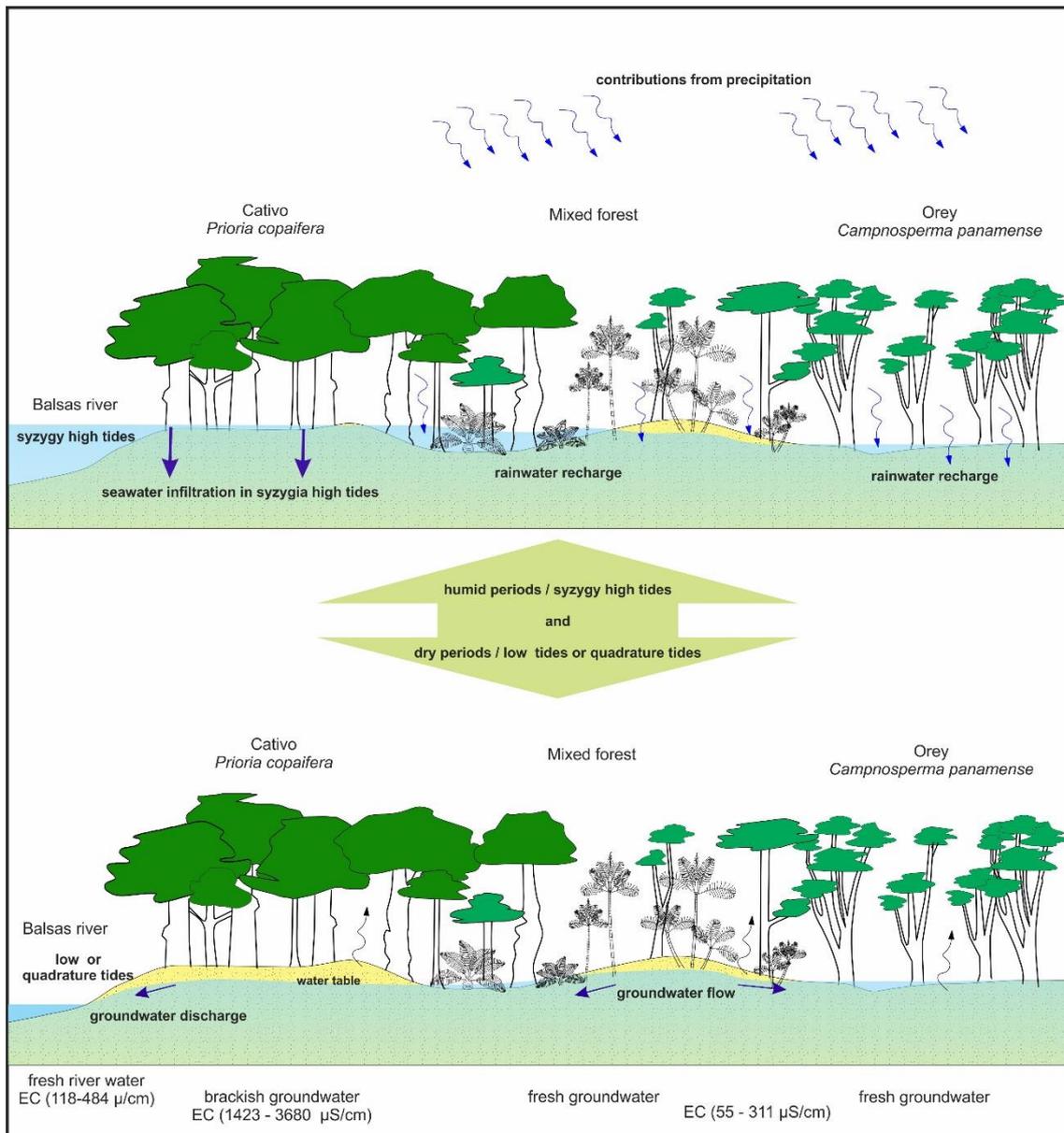
266 lower basin area, which shows up to 2 orders of magnitude more EC compared to the middle and
267 upper basin. On the other hand, in response to the influx of the tide in the middle and upper basin,
268 the levels rise because of the accumulation of river water that cannot drain into the estuary. These
269 sectors of the river, the middle and upper basin, accumulate fresh water. The tidal influence also
270 causes spatial variations in the major ion composition of the water (Fig. 5). The river presents
271 sodium-calcium bicarbonate and sodium chloride facies in the upper basin that transition to
272 sodium chloride in the middle and lower basin. Furthermore, the isotopic signal confirms the idea
273 that in the lower basin there is tidal influence, as samples from this area present greater isotopic
274 enrichment with trends towards seawater. In contrast, in the middle and upper basin, the signal is
275 like that of the local rainfall (Fig. 6).

276 The rises of the river caused by the syzygy high tides flood the wetland environments on the river
277 levee and the alluvial plain. In the middle river basin, the study of transect 1 allowed us to
278 recognize three wetland environments with different plant communities in which different
279 hydrological behaviors are recognized. Next to the river, on the levee, the cativo forest develops.
280 In this environment, the groundwater level sensors showed rises in accordance with the syzygy
281 high tides (Fig. 3), which indicates that the groundwater that supports this forest is recharged by
282 river water. The high tide floods the levee, and the river water infiltrates the sediments, causing
283 rapid rises in the water table. The groundwater in this environment of the transect 1 presents one
284 of the highest EC because of the tidal influence. It also exhibits brackish characteristics and
285 sodium chloride facies like those of the river in this basin sector. Isotopically, although a signal
286 like that of local rainfall is observed, there is a slight deviation towards the trend of waters affected
287 by mixing with tidal water (Fig. 6b).

288 In the mixed forest and in the ore forest located behind the river levee, the hydrodynamics and
289 water chemistry show a different behavior. The water table levels in these two wetland
290 environments do not vary with the tide. Only a series of slight increases in water table levels are
291 observed in the mixed forest during extraordinary spring tides (Fig. 3). In the water level records,
292 slight variations are observed, sometimes associated with rainfall. However, the water table is

293 very close to the surface or is outcropping, resulting in rejected recharge. Under these conditions,
294 a shallow water layer accumulates on the surface or drains through small streams that cross the
295 levee and discharge into the river. The electrical conductivity (EC) of the groundwater here is
296 very low in both environments (fresh groundwater), and hydrochemical facies may vary, although
297 sodium-calcium bicarbonate facies dominate (Fig. 5). The isotopic signal of groundwater and the
298 water that accumulates on the surface or drains through small streams is like that of rainwater. All
299 these hydrodynamic and hydrochemical characteristics indicate that these wetland forests are
300 primarily sustained by rainwater and do not receive tidal influence.

301 Laterally, the forest environments are connected through underground flow. However, these
302 lateral inputs are minimal compared to the flooding from tidal water in the cativo forest or the
303 contribution of rainfall in the mixed and ore forest. A summary of lateral hydrological flows
304 and connectivity is shown in Figure 7.



305

306 Figure 7: Diagram showing surface and groundwater flows and lateral hydrological connectivity
 307 in different wetlands environments in the middle basin of the Balsas River (Transect 1).

308

309 Transect 2 is in the lower basin, where tidal influence on the hydrodynamics and water chemistry
 310 of the river is more pronounced. In this wetland environment, the dominant vegetation is
 311 mangrove, with sectors identified as either dominated by caballero mangrove or salado mangrove.
 312 In both mangrove forests, the groundwater level increases during spring tide events (Fig. 4),
 313 indicating that the groundwater supporting the mangroves is primarily recharged by river water.

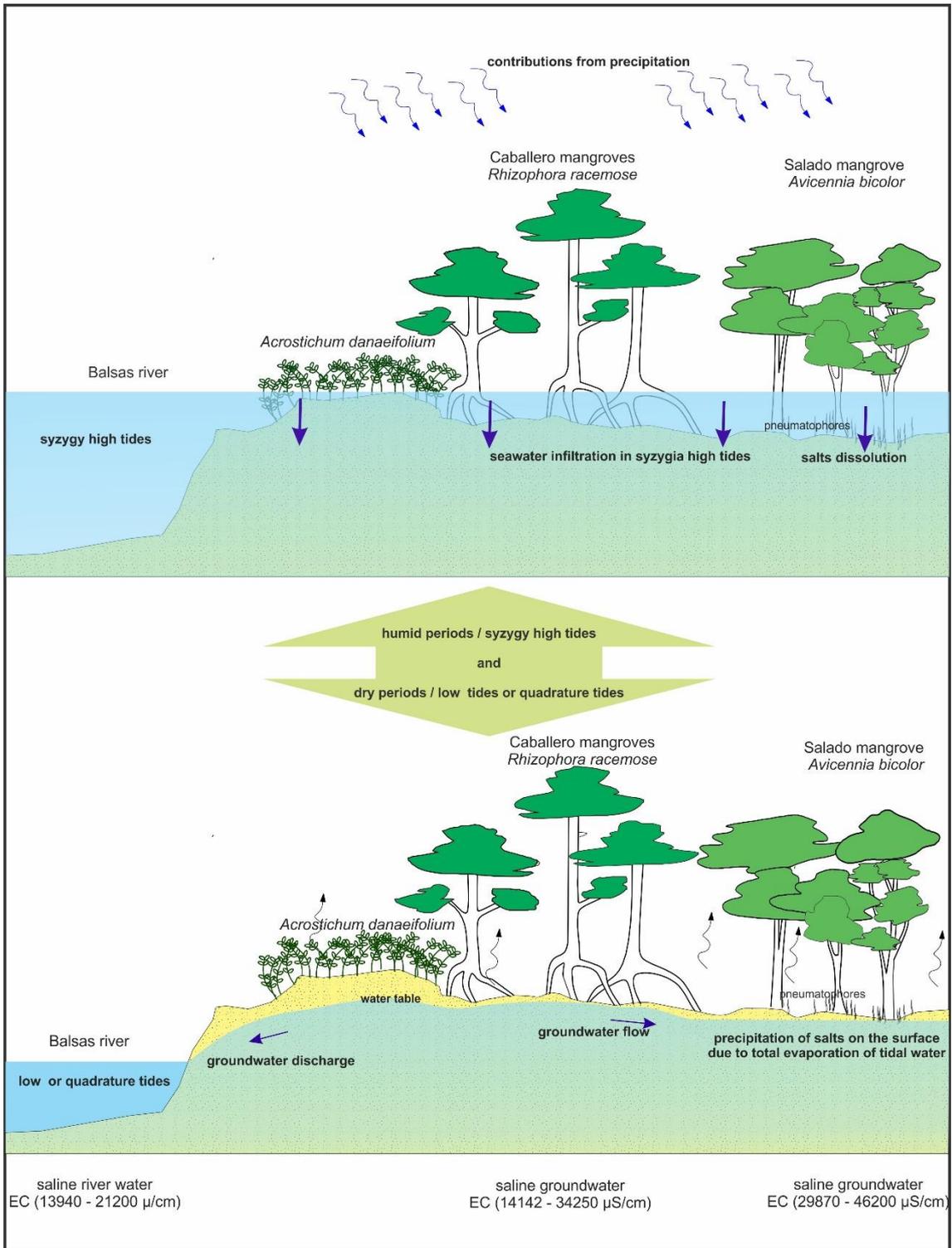
314 The high tide exceeds the river levee and floods the mangrove areas located behind it, with river
315 water infiltrating the sediments, causing rapid rises in the water table. This tidal influence causes
316 the groundwater to have high electrical conductivity, saline characteristics and sodium chloride
317 facies like those of the river in this basin sector.

318 In transect 2, the isotopic signal of groundwater shows a clear trend towards mixing with seawater
319 (Fig. 6a). The EC values in the mangrove groundwater are considerably higher than those in the
320 river water, resulting from salt dissolution processes (Fig. 6b) suggesting that the tidal water that
321 inundates the mangrove accumulates on the surface, and in some areas, it may completely
322 evaporate, forming saline precipitates on the substrate. These precipitates are later dissolved in
323 the next tidal inundation (Carol and Alvarez 2016; Galliari et al., 2021). High temperatures and
324 evapotranspiration rates contribute to the formation of evaporitic salts. This salinization process
325 is much more pronounced in the salado mangrove, which also experiences marked decreases in
326 the water table during periods of low precipitation and low spring tides (Fig. 4). While the
327 contribution of precipitation is not ruled out, the observed hydrodynamics and hydrochemistry
328 indicate that tidal influence is much more significant.

329 Laterally, the mangrove environments are connected to each other through underground flow.
330 However, these lateral contributions are minimal compared to the flooding caused by tidal water.
331 A summary of lateral hydrological flows and connectivity in the transect 2 area is shown in Figure
332 8.

333

334



335

336 Figure 8: Diagram showing surface and groundwater flow exchanges and lateral hydrological
 337 connectivity in different wetlands environments in the lower basin of the Balsas River (Transect
 338 2).

339

340 The study of transects 1 and 2 indicate that there are wetland environments that primarily depend
341 on tidal flow (wetlands in the lower basin and embankment sectors in the middle basin) and others
342 that rely mainly on rainfall (wetlands behind the embankment in depressed areas within the
343 middle basin). Similar behavior had been observed in studies conducted in wetland environments
344 associated with the Tuira River (Carol et al., 2020; 2021; 2022).

345 Hydrology plays a key role in the distribution and development of plant communities in wetlands
346 (Ridolfi et al., 2006; Dwire et al., 2006; Loheide and Gorelick, 2007; Muneeppeerakul et al., 2008).
347 However, despite observing similar hydrological behaviors between the Balsas and Tuira rivers'
348 wetlands, the observed plant communities are different. This seems to indicate that in
349 Matusagaratí, there are different plant communities adapted to similar hydrological conditions.

350 Ecohydrological models provide a useful tool for understanding interactions between surface
351 water, groundwater, and vegetation (Chui et al., 2011). In this regard, the obtained results and
352 conceptual models not only improve our understanding of the hydrological functioning of
353 different environments within the Matusagaratí wetland but also enable the identification of
354 vegetation adapted to each of the defined hydrological conditions. This constitutes a crucial tool
355 for developing water management guidelines and promoting sustainable management of wetland
356 areas, which, despite being legally protected, continue to face anthropogenic pressures from the
357 unsustainable development practices in the region.

358

359 **5. Conclusions**

360 This study identified the surface and groundwater flow exchanges and lateral hydrological
361 connectivity in different environments associated with the Balsas River in the Matusagaratí
362 Wetland. Each of these environments presents a characteristic type of vegetation, which shows
363 that hydrology conditions the development of the different plant species in the wetland.

364 The hydrodynamics of the Balsas River are significantly influenced by tides. In the lower basin,
365 tides enter through the estuary, causing rises in the river level during high tides. In the middle

366 basin, high tides lead to increases in river levels due to the accumulation of water that cannot
367 drain towards the estuary. The influx of tidal water causes variations in the salinity of the river
368 water. In the upper and middle basins, the water is freshwater with hydrochemical facies
369 characterized by bicarbonate-chloride composition and an isotopic signal like rainfall. On the
370 other hand, in the lower basin, the river water is saline with sodium chloride facies and isotopic
371 trends resembling seawater.

372 Spatial variations in hydrodynamic and hydrochemical behavior are also recognized in the
373 wetlands associated with the river. In the middle basin, environments such as cativo forests
374 develop on the levee, and mixed forests and orey forests develop in the alluvial plain. In these
375 areas, variations in water table levels and the major chemical and isotopic composition of
376 groundwater reveal different hydrological processes.

377 In the cativo forest, nearest the river basin, there is an interaction between wetland groundwater
378 and river water. This interaction takes place during spring tides when the Balsas River locally
379 floods the levee, and river water infiltrates, causing rises in the water table levels during these
380 tides. The input of river water results in brackish groundwater in the cativo forest with a slight
381 isotopic trend towards tidal water.

382 On the other hand, mixed forest and orey environments, developed in lower-lying alluvial plain,
383 exhibit hydrodynamics and hydrochemistry primarily associated with precipitation inputs. In
384 these environments, water levels do not fluctuate in relation to tides, and the water is freshwater
385 with an isotopic composition like rainfall. These wetland areas are drained by small streams where
386 surface water exhibits similar chemical and isotopic characteristics.

387 In the lower basin, the hydrodynamics and hydrochemistry of the Balsas River and the mangrove
388 wetlands, specifically mangle caballero and mangle salado, show a significant tidal influence.
389 Syzygy tides inundate these mangroves with saline water, and the water infiltrates, causing rises
390 in groundwater levels. The wetland's groundwater exhibits sodium chloride hydrochemical facies
391 and an isotopic signal associated with tidal contributions. However, groundwater in the

392 mangroves, especially in mangle salado, has higher electrical conductivity than river water due to
393 the dissolution of salts precipitated during the complete evaporation of tidal water during periods
394 when the wetland is not flooded.

395 The findings of this study on the Balsas River, along with previous work on the Tuira River,
396 highlight the complex hydrological dynamics of the Matusagaratí wetland. The data provided
397 forms a foundation for the management of this extensive tropical wetland, which, despite having
398 protection initiatives, could be hydrologically impacted by unsustainable practices, such as the
399 construction of drains and embankments for rice cultivation, occurring in the region.

400

401 **Acknowledgments**

402 This research would not have been possible without the generous funding from the National
403 Secretariat for Science, Technology, and Innovation (SENACYT) through the project Hydrology,
404 Carbon Reserves, Plants, and Fish Diversity of Matusagaratí (PFID-FID-2021-114). We also
405 express our gratitude to the Ministry of Environment for granting the necessary permits for the
406 collection and exportation of soil, water, and plant samples. Many thanks to our thesis student,
407 Ian Deago, for his willingness and endurance during the challenging fieldwork conditions. Finally,
408 our heartfelt gratitude to Hayro Cunampio, Jorge Tomi, Aurelio Flaco, Jhon Flaco, Eduardo
409 Garabato, Ismael Flaco y Antonio Martínez, residents of Afrodescendant and Emberá
410 communities, who always supported us with their expertise and knowledge of these valuable
411 wetland ecosystems.

412

413 **1. References**

414

415 Adame M. F., Franklin H., Waltham N. J., Rodriguez S., Kavehei E., Turschwell M. P., Ronan M.
416 2019. Nitrogen removal by tropical floodplain wetlands through denitrification. *Marine*
417 *and Freshwater Research*, 70 (11), 1513-1521.

418 Ahmed S., 2015. Local Level Perspectives of Wetland Management Policy and Practices in
419 Bangladesh: A Case of Hakaloki Haor (master's thesis). The University of Manitoba,
420 Winnipeg, Manitoba.

421 Aparicio K. 2021. Aves de Matusagaratí. Informe de resultados del Proyecto Hidrología,
422 Vegetación y Avifauna del Complejo de Humedales de Matusagaratí, Darién. 38 pp.

423 Barbier E.B., 2011. Wetlands as natural assets. *Hydrological Sciences Journal* 56, 1360-1373.

424 Candanedo I. 2021. Matusagaratí: el Pantanal de Panamá. Resumen para tomadores de decisión.
425 Universidad Tecnológica de Panamá. Universidad Tecnológica de Panamá. Secretaría
426 Nacional de Ciencia, Tecnología e Innovación (SENACYT). Panamá. 24 pp.

427 Carol E., Alvarez M.P. 2016. Processes regulating groundwater chloride content in marshes under
428 different environmental conditions: A comparative case study in Península Valdés and
429 Samborombón Bay, Argentina. *Continental Shelf Research*, 115, 33-43.

430 Carol E., Alvarez M.P., Candanedo I., Saavedra S., Arcia M., Franco A. 2020. Surface water-
431 groundwater interactions in the Matusagaratí wetland, Panama. *Wetlands, Ecology and*
432 *Management*. DOI.org/10.1007/s11273-020-09762-9.

433 Carol E., Alvarez M.P., Candanedo I., Arcia M. 2021. Estudiando el funcionamiento hidrológico
434 del Humedal de Matusagaratí. Universidad Tecnológica de Panamá. Secretaría Nacional
435 de Ciencia, Tecnología e Innovación (SENACYT). Panamá. 32 pp.

436 Carol E., Alvarez M.P., Santucci L., Candanedo I., Arcia M. 2022. Origin and dynamics of surface
437 water – groundwater flows that sustain the Matusagaratí Wetland, Panamá. *Aquatic*
438 *Sciences* 84: 16.

439 CREHO (Centro Regional Ramsar para la Capacitación e Investigación en Humedales en el
440 hemisferio occidental). 2015. Diagnóstico socioambiental, Laguna de Matusagaratí,
441 CREHO, CEASPA, ACD.

442 Chui T. F., Low S. Y., Liong S. Y. 2011. An ecohydrological model for studying groundwater–
443 vegetation interactions in wetlands. *Journal of Hydrology*, 409(1-2), 291-304.

444 Costanza R., de Groot R., Farber S., Grasso M., Hannon B., Limburg K., Van Den Belt, M. 1998.
445 The value of the world# s ecosystem services and natural capital. *Ecological economics*,
446 25 (1), 3-15.

447 Cuthbert R.N., Wasserman R.J., Keates C., Dalu T., 2022. Food webs. In: Dalu, T., Wasserman,
448 R.J. (Eds.), *Fundamentals of Tropical Freshwater Wetlands: From Ecology to*
449 *Conservation Management*. Elsevier, Cambridge.

450 De Groot R. S., Stuij M., Finlayson C., Davidson N. 2006 Valuing wet- lands: guidance for
451 valuing the benefits derived from wetland ecosystem services Ramsar Technical Report
452 No. 3/CBD Technical Series No. 27 Technical Series No. 27(Montreal: Ramsar
453 Convention Secretariat, Gland, Switzerland & Secretariat of the Convention on
454 Biological Diversity)

455 Dube T., Pinceel, T., De Necker L., Wepener V., Lemmens P., Brendonck L. 2019. Lateral
456 hydrological connectivity differentially affects the community characteristics of multiple
457 groups of aquatic invertebrates in tropical wetland pans in South Africa. *Freshwater*
458 *biology*, 64(12), 2189-2203.

459 Dwire K. A., Kauffman J. B., Baham J. E. 2006. Plant species distribution in relation to water-
460 table depth and soil redox potential in montane riparian meadows. *Wetlands*, 26(1), 131-
461 146.

462 Erwin K. L. 2009. Wetlands and global climate change: the role of wetland restoration in a
463 changing world *Wetlands Ecol.Manage.* 17 71

464 Fabrega J., Nakaegawa T., Pinzón R., Nakayama K., Arakawa O., Sousei T. 2013. Hydroclimate
465 projections for Panama in the late 21st Century. *Hydrological Research Letters*, 7(2), 23-
466 29.

467 Galliari J., Santucci L., Misseri L., Carol E., Alvarez M.P. 2021. Processes controlling
468 groundwater salinity in coastal wetlands of the southern edge of South America. *Science*
469 *of the Total Environment*, 754, 141951.

470 Gonfiantini, 1978. Standards for stable isotope measurements in natural compounds. *Nature* 271
471 (5645), 534.

472 Grauel W.T. 2004. Ecology and management of wetland forests dominated by *Prioria copaifera*
473 in Darien, Panama. PhD. Thesis. University of Florida. Florida, USA. 163 pp.

474 Ibáñez A., R Flores. 2020. *Phyllanthus fluitans* (Phyllanthaceae): a new record of an aquatic plant
475 for the flora of Panama. *Acta Botanica Mexicana* 128: e1767. DOI:
476 10.21289/abm128.2021.1767.

477 Ivory S. J., McGlue M. M., Spera S., Silva A., Bergier I. 2019. Vegetation, rainfall, and pulsing
478 hydrology in the Pantanal, the world's largest tropical wetland. *Environmental Research*
479 *Letters*, 14(12), 124017.

480 Kern Z., Harmon R.S., Fórizs I. 2016. Stable isotope signatures of seasonal precipitation on the
481 Pacific coast of central Panama. *Isotopes in Environmental and Health Studies* 52: 128-
482 140. doi: 10.1080/10256016.2015.1016021.

483 Loheide S. P., Gorelic, S. M. 2007. Riparian hydroecology: a coupled model of the observed
484 interactions between groundwater flow and meadow vegetation patterning. *Water*
485 *Resources Research*, 43(7).

486 López H., Cunampio H. 2023. Pesca Artesanal y Estrategia de Supervivencia en Comunidades de
487 Matusagarati, Darién. Proyecto Hidrología, Carbono, Plantas y Peces de Matusagaratí.
488 Informe de campo. 29 pp.

489 Ministerio de Ambiente. 2016. Estudio Técnico Justificativo para la creación del área protegida
490 Humedal Laguna de Matusagaratí. Dirección Nacional de Áreas Protegidas y Vida
491 Silvestre. 99 páginas.

492 Muneeppeerakul C. P., Miralles □ Wilhelm F., Tamea S., Rinaldo A., Rodríguez □ Iturbe I. (2008).
493 Coupled hydrologic and vegetation dynamics in wetland ecosystems. *Water Resources*
494 *Research*, 44(7).

495 Ortiz O.O., Ibáñez A., Trujillo-Trujillo E., Croat T.B. 2020. The emergent macrophyte
496 *Montrichardia linifera* (Arruda) Schott (Alismatales: Araceae), a rekindled old friend
497 from the Pacific Slope of lower Central America and western Colombia. *Nordic Journal*
498 *of Botany* 38(9): 1-10. <https://doi.org/10.1111/njb.02832>

499 Ortiz O.O., Croat T.B., Rodríguez-Reyes O., Ceballos J., Cedeño-Fonseca M., Mora M. 2022.
500 Taxonomic Novelties in *Philodendron* subg. *Philodendron* (Araceae) from Panama.
501 *Novon* 30: 18-42.

502 Ridolfi L., D'Odorico P., Laio F. 2006. Effect of vegetation-water table feedbacks on the stability
503 and resilience of plant ecosystems, *Water Resour. Res.*, 42, W01201,
504 doi:10.1029/2005WR004444.

505 Wasserman R. J., Dalu T. 2022. Tropical freshwater wetlands: an introduction. In *Fundamentals*
506 *of Tropical Freshwater Wetlands* (pp. 1-22). Elsevier.

507 Xiao K., Li H., Shananan M., Zhang X., Wang X., Zhang Y., Zhang X., Liu H. 2019. Coastal
508 water quality assessment and groundwater transport in a subtropical mangrove swamp in
509 Daya Bay, China. *Science of the Total Environment*, 646, 1419-1432.

Declaration of interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

Author Contributions

Carol E.: Investigation, field surveys, conceptualization, data curation, formal analysis, writing—original draft preparation, writing—review, and editing.

Alvarez M.P.: Investigation, field surveys, conceptualization, data curation, formal analysis, writing—original draft preparation, writing—review, and editing.

Arcia M.: Field surveys, data acquisition, data curation, writing—review, and editing.

Candanedo I.: Field surveys, project administration, funding acquisition.